

CO₂ Uptake of the Ocean - Parameter Optimization in Biogeochemical Models

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Outline

- Why study CO₂ uptake?
- How is it working?
- What are the models?
- Computing a steady seasonal cycle
- Full Jacobian Approach for optimization
- Brief look inside a biogeochemical model
- What is *Future Ocean*?

Why study CO₂ uptake?

CO₂ ...

- ... in the atmosphere: 800 Giga (10⁹) tons
pre-industrial: 600 Gigatons
- ... in the oceans: 40'000 Gigatons (atmosphere:ocean 1:50)



- ... exchange atmosphere-ocean: 90 Gigatons/year

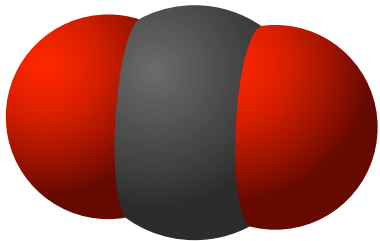


Image: Jacek FH GNU FDL

- today: 6 Gigatons emissions/year,
2 Gigatons are dissolved in the oceans
ocean decreases global warming, but: acidification
- questions: future CO₂ uptake?
how much remains in the ocean?
feedback effects: changes due to global warming?
sequestration opportunities, techniques, safety?
fertilization to increase uptake?

"Who" is taking up CO₂?

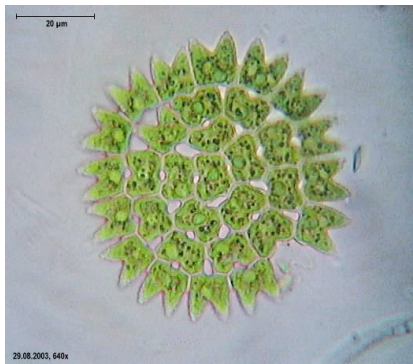
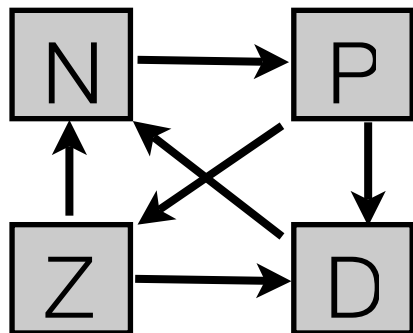


Bild: Dr. Ralf Wagner
GNU FDL



Image: MAR-ECO Øystein Paulsen GNU FDL



- "biological pump":
 - photosynthesis of phytoplankton
 - phytoplankton as nutrient of zooplankton
 - dead organic material sinking to the bottom, little mixing of surface and bottom layers
- "carbon pump":
 - sinking of calcium carbonate (inorganic)
- All that happens in the ocean circulation

Coupled Ocean Biogeochemistry

Ocean
circulation

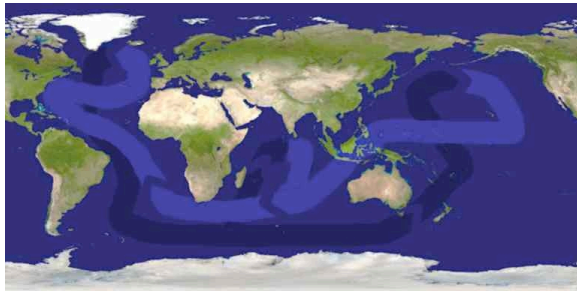
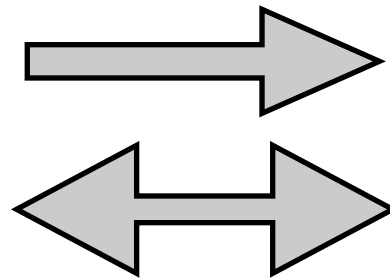


Bild: Nerd GNU FDL

First concept:

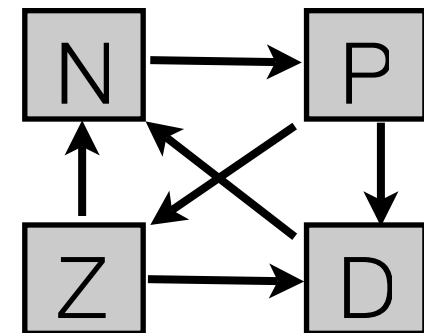
use
computed
circulation



Second concept:

full
interaction

Biogeo-
chemistry



Differences

Ocean circulation

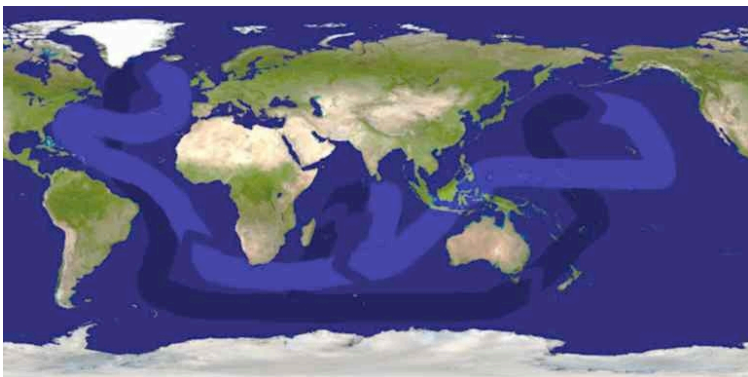
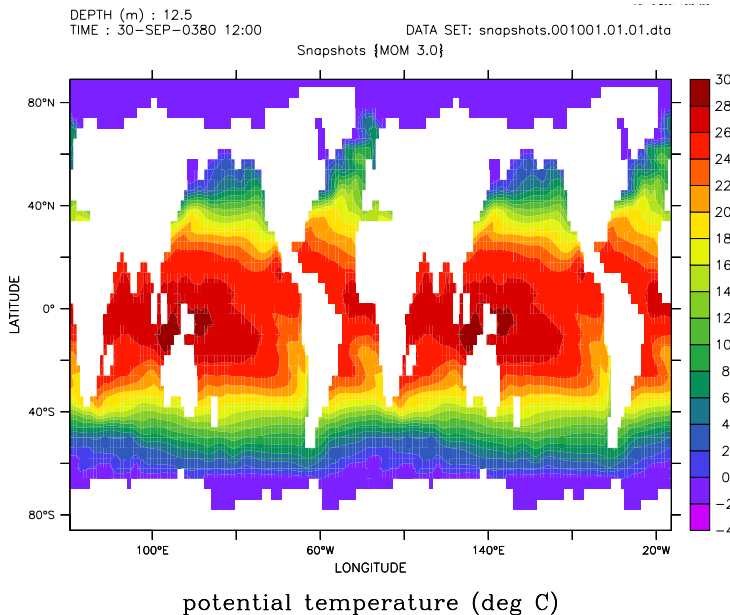


Bild: Nerd GNU FDL

- Equations are standard
- Some parts are focus of current research (e.g. diffusion)
- time-dependent, 3-D, nonlinear
- small scales: resolving turbulence would need about 10^{28} points
- thousands yeras needed to reach steady seasonal cycle
- Thus: Simplifications, modeling, model parameters

Ocean Circulation

- Ocean model (Navier-Stokes + tracer equations, hydrostatic + Boussinesq approximations)

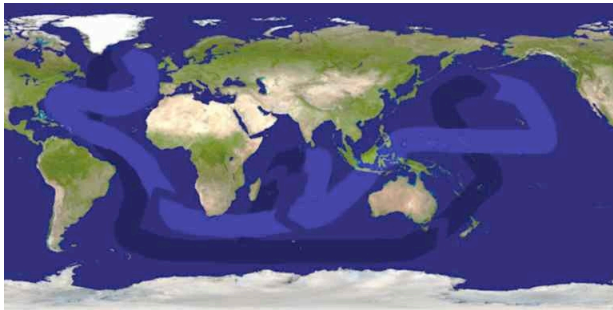


Image: Nerd GNU FDL

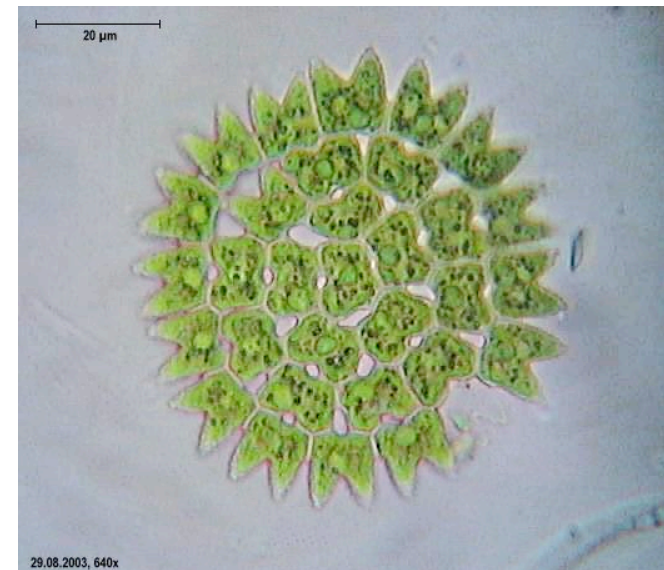
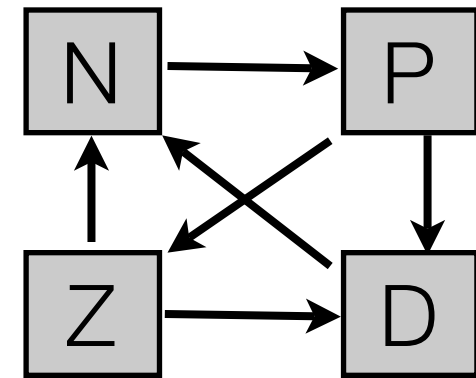
$$\begin{aligned}
 \vec{v}_t - \text{div}(\nu \nabla \vec{v}) + \vec{v} \cdot \nabla \vec{v} + \nabla p &= -2\vec{\Omega} \times \vec{v} - \rho \vec{g} \\
 \text{div} \vec{v} &= 0 \\
 S_t - \text{div}(\nu \nabla S) + \vec{v} \cdot \nabla S &= 0 \\
 T_t - \text{div}(\nu \nabla T) + \vec{v} \cdot \nabla T &= 0 \\
 \rho &= \rho(T, S)
 \end{aligned}$$

- numerical solution available (MITgcm, MOM, OPA ...)
- takes some 1000s years model time to become "steady", i.e. no change from year to year, i.e. periodic

Differences

- Transport equations, nonlinear coupling, easier compared to ocean model
- Not clear how many tracers to include
 - NPZD model: nutrients, phytoplankton, zooplankton, detritus
 - more tracers-> more equations-> more parameters -> but more information?
 - Modeling is current research topic

Biogeochemistry



... and Biogeochemistry

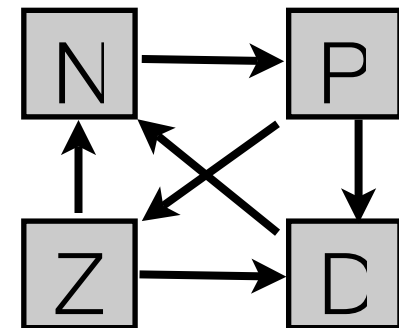
- Linear transport/advection-diffusion equations with nonlinear coupling/forcing:

$$\frac{\partial c_i}{\partial t} = \underbrace{\text{div}(\kappa \nabla c_i)}_{\text{diffusion}} - \underbrace{\text{div}(\vec{v} c_i)}_{\text{advection}} + \underbrace{q_i(c_1, \dots, c_N, T, S, p)}_{\text{tracer coupling}}$$

- for tracers, e.g. nutrients, phyto-, zooplankton, detritus

- Different numbers of tracers, equations, many parameters

- "more tracers -> better results" is not naturally true



- local (in gridpoints, except for sinking) biochemical processes in source terms q
- given "steady" circulation: separated tracer computation into a "steady" state, starting from uniform distribution, again 1000s of years model time

The seasonal cycle

- Using pre-computed velocity, temperature, salinity the biogeochemical equations

$$\frac{\partial c_i}{\partial t} = \text{div}(\kappa \nabla c_i) - \text{div}(\vec{v} c_i) + q_i(c_1, \dots, c_N, T, S, p)$$

- can be written as $\frac{dc_i}{dt} = A(t)c_i(t) + q(c(t), t)$

- where $A(t)$ can be pre-computed as *Transition Matrix* [Khatiwala et. al. '04]

- Time discretization gives for the vectors of tracers: $c^{k+1} = \Phi_k(c^k)$

- Looking for a periodic solution (steady seasonal cycle):

$$c^{k+n} = \underbrace{\Phi_{k+n-1} \circ \dots \circ \Phi_{k+1} \circ \Phi_k}_{=: \Phi}(c^k)$$

- i.e. a fixpoint of Φ

Computing the steady seasonal cycle

- Pseudo time-stepping (this is done by now): Compute until for some k (n fixed):

$$c^{k+n} = \Phi(c^k) = c^k$$

- vs. Newton's method for the fixpoint equation:

$$\Phi(c) = c \iff \Phi(c) - c = 0$$

- Requires Jacobian $\left(\frac{\partial \Phi}{\partial c} - I \right)$

- where the Jacobian of Φ is mainly given by the matrices A_k and the q'_k

- remember:

$$c^{k+n} = \underbrace{\Phi_{k+n-1} \circ \dots \circ \Phi_{k+1} \circ \Phi_k}_{=:\Phi}(c^k) \qquad \frac{dc_i}{dt} = A(t)c_i(t) + q(c(t), t)$$

Inexact Newton & Preconditioning

- The linear system in the Newton steps can be solved inexactly/iteratively (GMRES etc.) and needs only matrix-vector products with A_k q'_k
- Two nested iterations, only useful if it needs less matrix-vector multiplications than the pseudo time-stepping:
Newton * # linear iterations \leftrightarrow # pseudo time-steps

- Preconditioning:
$$\left(\frac{\partial \Phi}{\partial c} - I \right)^{-1} \approx - \sum_{k=0}^K \left(\frac{\partial \Phi}{\partial c} \right)^k$$

- again reduces to multiplication with A_k and q'_k
- Use bigger timesteps for the biogeochemics for the preconditioner

Computing a derivative for optimization

- Looking for a steady seasonal cycle/fixpoint, depending on parameters p :

$$\Phi(c(p), p) = c(p)$$

- Taking the derivative w.r.t to the parameter vector:

$$\frac{\partial \Phi}{\partial c} c'(p) + \frac{\partial \Phi}{\partial p} = c'(p) \iff \left(\frac{\partial \Phi}{\partial c} - I \right) c'(p) = -\frac{\partial \Phi}{\partial p}.$$

- Same system matrix as for the computation of the steady seasonal cycle.
- To be tested against classical adjoint technique or oneshot method for pseudo-timestepping (see Adel Hamdi's talk)

Looking inside a biogeochemical model

- Oschlies, Garçon 1999: $c=(N,P,Z,D,DOM)$

$$q(c) = \begin{pmatrix} 0 & 0 & \gamma_2 & \mu_4 & \mu_5 \\ 0 & -\mu_2 & 0 & 0 & 0 \\ 0 & 0 & -\mu_2 & 0 & 0 \\ 0 & (1-\sigma)\mu_2 & 0 & -\mu_4 & 0 \\ 0 & \sigma\mu_2 & 0 & 0 & -\mu_5 \end{pmatrix} c + \begin{pmatrix} -J(c_1, c_2)c_1 \\ J(c_1, c_2)c_1 - G(c_2)c_3 \\ \gamma G(c_2)c_3 - \mu c_3^2 \\ (1-\sigma)((1-\gamma_1)G(c_2)c_3 + \mu_3 c_3^2) \\ \sigma((1-\gamma_1)G(c_2)c_3 + \mu_3 c_3^2) \end{pmatrix} + \begin{pmatrix} 0 \\ 0 \\ 0 \\ -\omega \frac{\partial c_4}{\partial z} \\ 0 \end{pmatrix}$$

$$G(c_2) = \frac{g\epsilon c_2^2}{g + \epsilon c_2^2}$$

Summary

- CO₂ uptake is a crucial point in climate research
- One-way coupling between ocean circulation + biogeochemistry allows to precompute circulation
- Newton to obtain steady seasonal cycle, exploiting structure of Jacobian
- Optimization: up to now: GA, finite differences, aim: Full Jacobian approach + Algorithmic Differentiation of biogeochemical part, one-shot approach
- Analysis of model equations w.r.t. existence, uniqueness
- Work in progress ...

Ozean der Zukunft

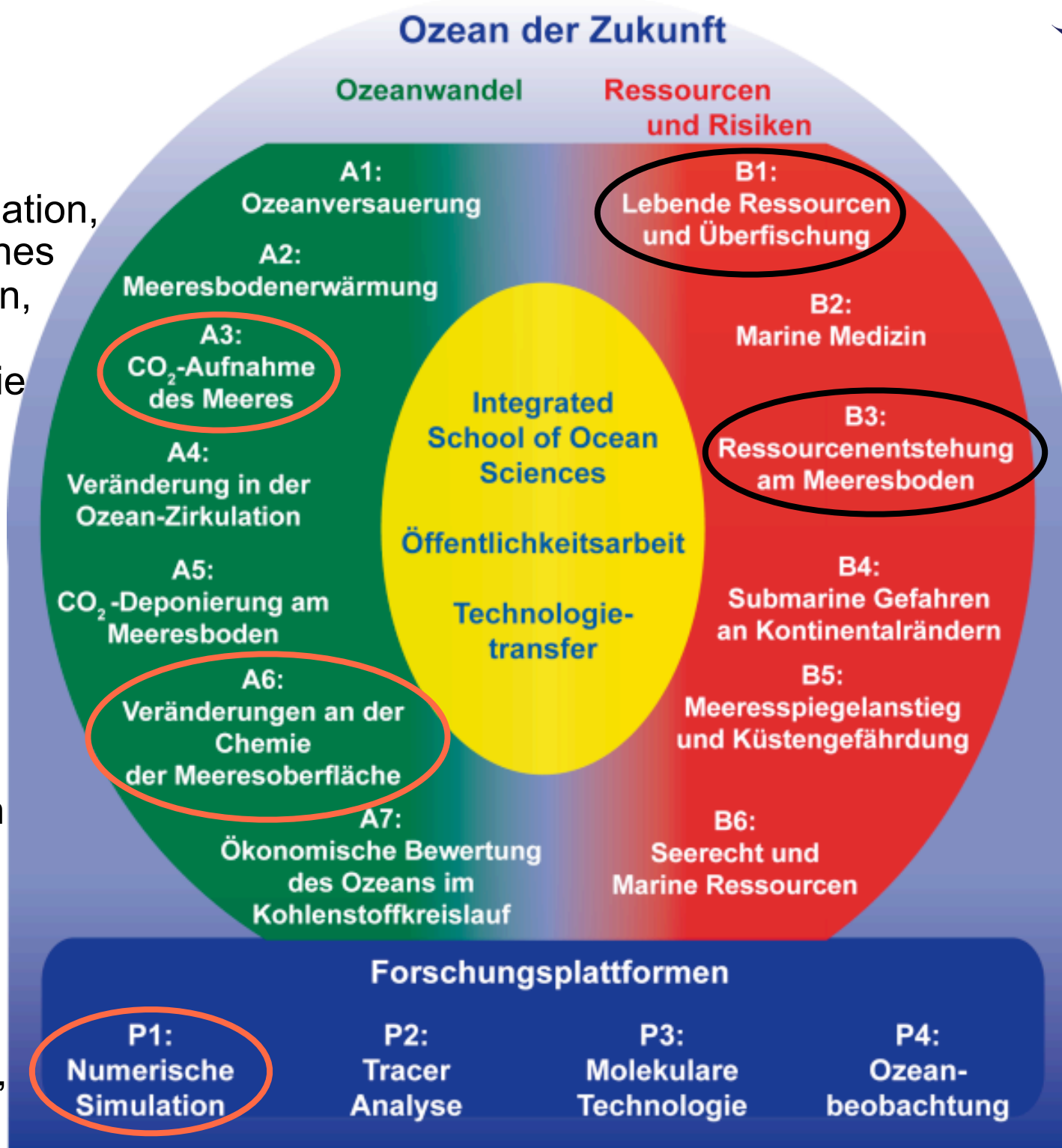
Parameter-
optimierung,
Datenassimilation,
Algorithmisches
Differenzieren,
Ozean +
Biogeochemie

Optimierung/
Steuerung bei
ökonomischen
Modellen (ODEs)

Parameter-
identifikation

Inverse
Modellierung
Stokes +
nicht-
Newtonische
Fluide

Numerik,
Optimierung,
Kopplung

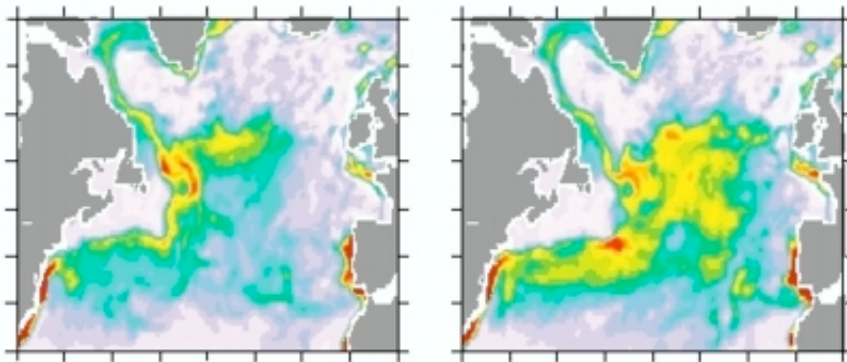


The Christian Albrechts University Kiel
invites to the

4. Scientific Computing Seminar

with special emphasis on

Parameter Estimation, Optimization and Inverse Modeling in Geosciences



16.-18.06.2008, Kiel



Workshop Program

Keynote Speakers

The following experts confirmed to give survey talks of 50 minutes:

- Dr. Peter Bayer (ETH Zürich)
- Prof. Roland Becker (Université de Pau et des Pays de l'Adour)
- Prof. Michael Hinze (Uni Hamburg)
- Prof. Rupert Klein (ZIB Berlin und PKI Potsdam)
- Dr. Samar Khatiwala (New York)
- Dr. Michael Vossbeck (FastOPT, Hamburg)
- Prof. Andrea Walther (TU Dresden)

Contributed talks - **Call for Papers**

We offer the possibility to deliver 30 minutes talks. Please send an abstract within 11 April 2008 via the registration link.

Thank you!!!